

# ARTICLE

Received 23 Jul 2013 | Accepted 2 Jun 2014 | Published 26 Jun 2014

DOI: 10.1038/ncomms5270

# Differentiating moss from higher plants is critical in studying the carbon cycle of the boreal biome

Wenping Yuan<sup>1,2,\*</sup>, Shuguang Liu<sup>3,\*</sup>, Wenjie Dong<sup>1,\*</sup>, Shunlin Liang<sup>4,5</sup>, Shuqing Zhao<sup>6</sup>, Jingming Chen<sup>7</sup>, Wenfang Xu<sup>1</sup>, Xianglan Li<sup>4</sup>, Alan Barr<sup>8</sup>, T. Andrew Black<sup>9</sup>, Wende Yan<sup>10</sup>, Mike L. Goulden<sup>11</sup>, Liisa Kulmala<sup>12</sup>, Anders Lindroth<sup>13</sup>, Hank A. Margolis<sup>14</sup>, Yojiro Matsuura<sup>15</sup>, Eddy Moors<sup>16</sup>, Michiel van der Molen<sup>17</sup>, Takeshi Ohta<sup>18</sup>, Kim Pilegaard<sup>19</sup>, Andrej Varlagin<sup>20</sup> & Timo Vesala<sup>21</sup>

The satellite-derived normalized difference vegetation index (NDVI), which is used for estimating gross primary production (GPP), often includes contributions from both mosses and vascular plants in boreal ecosystems. For the same NDVI, moss can generate only about one-third of the GPP that vascular plants can because of its much lower photosynthetic capacity. Here, based on eddy covariance measurements, we show that the difference in photosynthetic capacity between these two plant functional types has never been explicitly included when estimating regional GPP in the boreal region, resulting in a substantial overestimation. The magnitude of this overestimation could have important implications regarding a change from a current carbon sink to a carbon source in the boreal region. Moss abundance, associated with ecosystem disturbances, needs to be mapped and incorporated into GPP estimates in order to adequately assess the role of the boreal region in the global carbon cycle.

<sup>1</sup> State Key Laboratory of Earth Surface Processes and Resource Ecology, Beijing Normal University, Beijing 100875/Zhuhai 519087, China. <sup>2</sup> State Key Laboratory of Cryospheric Sciences, Cold and Arid Regions Environmental and Engineering Research Institute, Chinese Academy of Sciences, Lanzhou, Gansu 730000, China. <sup>3</sup> U.S. Geological Survey, Earth Resources Observation and Science Center, Sioux Falls, South Dakota 57198, USA. <sup>4</sup> State Key Laboratory of Remote Sensing Science, Jointly Sponsored by Beijing Normal University and Institute of Remote Sensing Applications, Chinese Academy of Sciences, College of Global Change and Earth System Science, Beijing 100875, China. <sup>5</sup> Department of Geography, University of Maryland, College Park, Maryland 20742, USA. <sup>6</sup> College of Urban and Environmental Sciences, Peking University, Beijing 100871, China. <sup>7</sup> Department of Geography, University of Toronto, Toronto, Ontario, Canada M5S 3G3. <sup>8</sup> Climate Research Division, Environment Canada, Saskatoon, Saskatchewan, Canada S7N 3H5. <sup>9</sup> University of British Columbia, Vancouver, British Columbia, Canada V6T 1Z4.<sup>10</sup> State Engineering Laboratory of Southern Forestry Applied Ecology and Technology, Central South University of Forestry and Technology, Changsha, Hunan 410004, China.<sup>11</sup> Department of Earth System Science, University of California, Irvine, California 92697-3100, USA. <sup>12</sup> Department of Forest Sciences, P.O. Box 27, FI-00014 University of Helsinki, Finland. <sup>13</sup> Department of Earth and Ecosystem Sciences, Lund University, Sölvegatan 12 223 62, Sweden. <sup>14</sup> Centre d'Étude de la Forêt, Faculté de foresterie, de géographie et de géomatique, Université Laval, Québec, Canada G1V 0A6.<sup>15</sup> International Forest Informatics Laboratory, Forestry and Forest Products Research Institute, Ibaraki 305-8687, Japan.<sup>16</sup> Wageningen UR, Alterra, Earth System Science and Climate Change Group, PO Box 47, 6700 AA Wageningen, The Netherlands. <sup>17</sup> Department of Hydrology and Geo-Environmental Sciences, Vrije Universiteit Amsterdam, 1081 HV Amsterdam, The Netherlands. <sup>18</sup> Graduate School of Bioagricultural Sciences, Nagoya University, Furo-cho, Chikusa-ku, Nagoya 464-8601, Japan.<sup>19</sup> Department of Chemical and Biochemical Engineering, Technical University of Denmark, DK-2800 Kongens Lyngby, Denmark.<sup>20</sup> Sukachev Laboratory of Biocenology, Severtsov Institute of Ecology and Evolution RAS, Leninsky pr.33, Moscow 119071, Russia.<sup>21</sup> Department of Physics, P.O. Box 48, FI-00014, University of Helsinki, Finland. \* These authors contributed equally to this work. Correspondence and requests for materials should be addressed to W.Y. (email: yuanwpcn@126.com) or to S.L. (email: sliu@usgs.gov) or to W.D. (email: dongwj@bnu.edu.cn).

he boreal biome accounts for 7% of global plant biomass and 50% of the global belowground organic carbon pool<sup>1,2</sup>. A recent study suggests that boreal forests in the northern high latitudes have sequestered about 0.5 PgC per year since 1990, which represents a significant portion of the estimated  $\sim 1.0$  PgC per year global net forest sink<sup>3</sup>. Moreover, climatic changes associated with increasing atmospheric CO<sub>2</sub> may be greatest at high latitudes<sup>4</sup>, and these changes may have important ramifications for the boreal biome, which may in turn influence the rate of regional and global climatic changes<sup>5</sup>. Therefore, understanding and quantifying the spatial and temporal changes of carbon fluxes over the boreal biome is a high priority.

Vegetation gross primary production (GPP) is the largest terrestrial carbon flux and one of the major components controlling land-atmosphere CO<sub>2</sub> exchange. A small error in simulating vegetation production can substantially alter the magnitude of the estimated global carbon balance<sup>6</sup>. Previous studies showed wide and inconclusive results regarding vegetation production over the high-latitude boreal biome. Globally, a comparison of 15 models showed the largest model differences in the boreal region (that is, boreal woodlands and boreal forests)<sup>7</sup>. Model estimates of GPP for boreal North America varied between 2.6 and 11.6 Pg C per year, a larger range than that of temperate North America<sup>8</sup>. Estimates of net primary production (NPP) in Canada's boreal forest also showed a large uncertainty, varying from 140 to  $444 \text{ g C m}^{-2}$  per year from three studies<sup>9–11</sup>.

Independently or as part of integrated ecosystem models, the satellite-based light use efficiency (LUE) approach has been used to estimate GPP and NPP at various spatial and temporal scales (see Methods). A recent study showed that an LUE model (the EC-LUE model) substantially overestimates vegetation primary production at young stands because it neglects the impacts of mosses<sup>12</sup>. Mosses often dominate the ground cover and contribute substantially to the satellite signals from young stands<sup>13</sup>. On the other hand, mosses with low LUE are fundamentally different from vascular plants with respect to

vegetation production<sup>14</sup>. However, almost all satellite-based models do not explicitly consider the impact of disturbance regimes on vegetation composition of the northern ecosystems, which may cause errors in the application of models using remotely sensed data for the boreal biome.

High-quality GPP estimates (using the semi-empirical fluxpartitioning algorithm) from the global network of eddy covariance (EC) flux towers make it possible to test the adequacy of algorithms for estimating GPP over large areas. In this study, we test the performance of three LUE models for predicting GPP (EC-LUE, MODIS-GPP and CASA)<sup>12,15,16</sup> at 26 EC flux towers in the boreal regions of North America, Europe and Asia (Supplementary Table 1; Supplementary Note 1), and quantify the relative contribution of different vegetation compositions to ecosystem GPP. The results show that three LUE models overestimate GPP at most sites, especially in younger stands. By separating the LUE and the relative contributions of mosses and vascular plants to the satellite signal using age information, the revised EC-LUE model successfully predicts GPP at all sites. Our regional results show that GPP has been substantially overestimated for the Canada boreal biome by the three LUE models that do not consider the impacts of moss. These results highlight the need to account for the impacts of moss on GPP across the boreal biome in models of terrestrial carbon dynamics.

#### Results

**Overestimated GPP over the boreal biome**. EC-LUE, MODIS-GPP and CASA explained about 68, 42 and 43% of the variation, respectively, of site-averaged GPP over their study periods across all sites (Supplementary Fig. 1a,c,d; see Methods). The relative predictive errors (RPEs, see Methods) from the three models decreased with stand age, and prediction errors approached zero for forests older than 50 years (Fig. 1a,c,d). All three models overestimated GPP at young forest sites. For example, in forests younger than 50 years, averaged RPE of the EC-LUE,



Figure 1 | The RPE of LUE models versus stand age. (a) The correlation relationship between stand age and RPE of EC-LUE, (b) EC-LUE-2p, (c) MODIS-GPP and (d) CASA models at all sites. Numbers within the graphs are site IDs (Supplementary Table 1).

MODIS-GPP and CASA models were 58, 61 and 67%, respectively (Fig. 1a,c,d). However, the three models offered accurate predictions for the old forest sites where the RPE was lower than 30% (Fig. 1a,c,d).

**Differentiation of moss from vascular plants**. We hypothesize that the overestimation of GPP, especially at young stands, was caused by the presence of moss and the lack of distinction between moss and higher plants in the satellite-derived normalized difference vegetation index (NDVI) data. Moss, one of the major groups of bryophytes, is a ubiquitous and dominant component of ground-level vegetation in the boreal biome. Moss is typically among the first ground layer species to become established after disturbance<sup>17</sup>. For example, at a study site in Interior Alaska, 70% of the surface was covered by moss, and only 30% of the surface was covered by grasses and deciduous shrubs 3 years after fire<sup>18</sup>.

Boreal forests are relatively open due to narrow tree crowns and low stand density of dominant trees; as a result, moss can substantially contribute to measured satellite signals<sup>19</sup>. This contribution from moss can be illustrated using a chronosequence of seven adjacent sites recovering from fire disturbances<sup>20</sup>. Figure 2 shows the leaf area index (LAI) of the



Figure 2 | Comparison of satellite NDVI field-measured LAI of vascular plants. (a) Comparison of field-measured LAI of vascular plants at seven adjacent sites in Canada; (b) comparison of MODIS NDVI at these seven sites. The sites (and ages in 2002) were CA-NS7 (4 years), CA-NS95 (7 years), CA-NS6 (13 years), CA-NS5 (21 years), CA-NS3 (38 years), CA-NS2 (72 years) and CA-NS1 (152 years). LAI of vascular plants was measured before the end of the growing season in 2004. NDVI values represent the mean NDVI values from May to September during 2002-2006, and the error bars indicate s.d. (see Methods).

vascular plants measured from the ground at these sites and their NDVI values derived from satellite data. NDVI includes the greenness contribution from both vascular plants and moss (see Methods). Measured LAI of vascular plants from the ground varied from 0.3 to  $7.4 \,\mathrm{m^2 m^{-2}}$  across these sites (Fig. 2a). In contrast, the NDVI values varied around 0.7, showing small differences across these seven sites (Fig. 2b), suggesting all these forests had a closed canopy<sup>21</sup>. The low LAI values in the young stands clearly indicated that their canopies were open. The discrepancy between field-based LAI measurements of vascular plants and satellite-derived NDVI suggests substantial contribution of moss to the satellite signal NDVI in the young stands.

Field experiments indicate that the LUE of moss is only about 10 to 50% of that of vascular plants<sup>22</sup>. To our knowledge, the lower photosynthetic capacity of moss has never been explicitly considered in any of the previous efforts in estimating GPP in the boreal biome. The importance of treating moss and vascular plants separately in estimating GPP can be demonstrated at site and regional scales using a modified EC-LUE model. We revised the EC-LUE model to explicitly include two contributions to ecosystem GPP: one from moss and the other from vascular plants (referred to as the EC-LUE-2p model hereafter). Intrinsic potential LUE values were specified for moss and vascular plants. In addition, we used K NDVI (see Methods) to represent the fractional contribution of moss to the NDVI signal at the ecosystem level. K\_NDVI, the only undetermined parameter of the EC-LUE-2p model, was inverted at the 26 flux tower sites against GPP estimates. Results show that after splitting the EC-LUE model into two parts, the GPP estimates were improved at the 26 flux tower locations and the mean absolute RPE decreased from 32 to 11% across these sites (Supplementary Figs 1, 2, 3 and 4). We found that the contribution of moss to the satellite greenness signal (K\_NDVI) was large in newly disturbed young forests, but the signal decreased exponentially with forest age (Fig. 3).

Independent field measurements of moss abundance also support the conclusion that the contribution of moss to the satellite greenness signal (K\_NDVI) decreases as forests age. A previous study along a wildfire chronosequence found that the



**Figure 3 | The contribution of moss to NDVI versus stand age.** K\_NDVI indicates the contribution of moss to the NDVI signal. Numbers in the graph are site IDs (Supplementary Table 1). The solid line and equation indicate the regression relationship between stand age and K\_NDVI. The two dashed lines indicate the 95% confidence band, and the regression equations are  $y = -0.11\ln(x) + 0.84$  (high) and  $y = -0.106\ln(x) + 0.64$  (low).



**Figure 4 | Regional GPP estimates of LUE models. (a)** The lowest, **(b)** the highest and **(c)** mean values of GPP estimates considering the impacts of moss based on the EC-LUE-2p model. **(d)** The difference in GPP simulations between the original EC-LUE and the EC-LUE-2p models (mean values). The low and high estimates represent the EC-LUE-2p estimates with 95% confidence. The original EC-LUE model does not consider the impacts of moss. The unit of GPP is  $g C m^{-2}$  per year.

bryophyte total cover was inversely correlated with forest age with higher cover in young stands<sup>23</sup>. In Alaskan forests, moss cover showed a unimodal distribution with time since fire occurrence, peaking 30–70 years post fire at the permafrost sites and 5–15 years at the permafrost-free sites<sup>24</sup>. The moss contribution to the NDVI signal is similar to the reported trend, which provides independent validation of the K\_NDVI derived from this study. Moreover, the partitioning of GPP between moss and vascular plants by the EC-LUE-2p model matched well with the observations at seven fire chronosequence sites<sup>20</sup> (Supplementary Fig. 5).

**Impacts on regional GPP estimations**. Using the improved EC-LUE-2p model (see Methods), the relationship between the contribution of moss to the NDVI signal and forest age derived from 26 flux tower locations (Fig. 3), and the spatial distribution map of forest  $age^{25}$ , we estimated GPP to be 2.44 Pg C per year for the Canadian boreal biome (Fig. 4c; see Methods). Moreover, we quantified the uncertainty of regional GPP estimations resulting from the relationship between stand age and K\_NDVI, shown in Fig. 3 (see Methods). The EC-LUE-2p GPP estimates varied from 2.15 to 2.67 Pg C per year over the entire study area for a 95% confidence limit (Fig. 4a,b).

In comparison, EC-LUE, MODIS-GPP and CASA generated estimates of 3.30, 3.69 and 3.98 PgC per year, respectively, for the same region when moss contribution was not explicitly considered. This finding represents an overestimation of 34–63% of the result from the EC-LUE-2p model (Fig. 4d). The EC-LUE-2p-simulated average GPP was  $388 \text{ g Cm}^{-2}$  per year; even the maximum estimation of GPP was  $425 \text{ g Cm}^{-2}$  per year, substantially lower than the  $601 \text{ g Cm}^{-2}$  per year recently obtained over the global boreal biome by diagnostic models<sup>26</sup>.

This result clearly demonstrates the importance of moss, disturbance and forest development stage in estimating vegetation production in the boreal biome.

#### Discussion

GPP is a critical component of the carbon cycle because it signifies carbon uptake by ecosystems. GPP estimates have been directly or indirectly used to evaluate regional carbon budgets<sup>27</sup>. For example, GPP maps were used to provide prior information for atmospheric inversion experiments<sup>28</sup>, and studies have shown inverse estimates of CO<sub>2</sub> fluxes in boreal Asia to be very sensitive to these prior estimates<sup>29</sup>. Unfortunately, spatial patterns of the contributions from moss and/or vascular plants to satellite signals have never been separately mapped and used in any of the previous efforts to estimate regional GPP. Consequently, current estimates of regional GPP are likely much higher than reality, as the photosynthetic capacity of moss is implicitly and artificially elevated to the same level as that of vascular plants.

The northern latitude boreal biome is generally considered to be acting as a considerable sink for atmospheric  $CO_2$  based on the results from both 'top-down' and 'bottom-up' approaches<sup>30</sup>. However, the precise role that the northern high-latitude region plays in the global carbon budget is not well understood because of model uncertainties<sup>31</sup>. For example, a study based on a nested inverse modelling system indicates Canada's sink to be large ( $0.34 \pm 0.14$  Pg C per year)<sup>28</sup>. However, ecosystem modelling and inventory-based analysis found Canada's forest to be a weak sink of 0.05-0.07 Pg C per year<sup>11,25</sup>. Regional carbon balance strongly depends on the accuracy of estimated GPP. An overestimation of regional GPP, as we have found in this study, has important ramifications for the carbon cycle in the boreal region. Although we only examined LUE models in this study, the conclusion is applicable to GPP models based on enzyme kinetics. A recent study reported large differences in enzyme kinetics between moss and vascular plants<sup>32</sup>. Moreover, the satellite-based LAI was used to drive physiological models. For example, GPP simulations from the Breathing Earth System Simulator, which integrates the MODIS LAI product with the enzyme kinetics model to estimate GPP, were most sensitive to LAI<sup>33</sup>. However, these physiological models did not differentiate the different contributions of moss and higher plants to satellite signal and ecosystem GPP. Therefore, this study highlights the importance of incorporating distinct structural and functional properties of ecosystems into regional modelling frameworks.

The implication of the results from this study to the total carbon budget of the boreal biome is important but quantitatively unknown as carbon sinks and sources are determined by both GPP and respiration. Given that plant autotrophic respiration (Ra) is proportional to GPP (about half)<sup>34</sup> for both moss and vascular plants<sup>35</sup>, previous Ra estimates are likely proportionally overestimated. However, Ra overestimates could not offset all GPP overestimates as the proportionality is about half, and the remaining half would result in overestimates of NPP (GPP minus Ra). The impacts of failing to account for the presence of moss on heterotrophic respiration are difficult to evaluate because heterotrophic respiration is affected by many processes, including NPP allocation among moss and higher plants and mortality; we do not yet have a model that comprehensively represents all these processes mechanistically. Nevertheless, some rough estimation can illustrate the consequence of GPP overestimates on the regional carbon budget. Assuming that the overestimated NPP is half of the overestimated GPP and is either completely returned to the atmosphere via heterotrophic respiration (low impact of GPP overestimation) or entirely accumulated in terrestrial ecosystems (high impact of GPP overestimation), the net impact on net carbon gain would range from 0 to 0.43 (EC-LUE), 0 to 0.62 (MODIS-GPP) and 0 to 0.77 Pg C per year (CASA) according to the GPP differences between EC-LUE-2p and these models. It is not likely that all overestimated NPP using the conventional approaches would be respired via heterotrophic respiration or completely retained by the ecosystems; the correct answer probably falls in the middle. If these considerations were incorporated in previous carbon sink estimates<sup>25,28</sup>, Canadian forests could be a source rather than a sink. In summary, the impacts of failing to differentiate the contributions of mosses and higher plants to the carbon cycle are substantial. Future research on the impact of moss on the carbon cycle, especially ecosystem respiration, is needed to reduce the uncertainty in estimating the regional carbon sink strength.

Our results suggest that characterizing plant composition and percent cover of bryophytes in the boreal biome is critical for quantifying the impacts of land surface dynamics on climate change because the boreal biome is the second largest biome in the world<sup>36</sup>. Bryophytes play an important role in ecosystem structure and function of boreal forests and have a profound impact on terrestrial carbon, water and nutrient budgets<sup>37</sup>. The importance of bryophytes and the lack of capability in mapping their dynamics thus highlight the need to develop remote sensing approaches to quantify the composition and percent cover of bryophytes in the region<sup>13</sup>. Even though MODIS data are gridded, the sensor flies over a given point on the ground at different angles over time. Therefore, in theory, the contribution of K\_NDVI should be different depending on the off-nadir view angle. This would change the MODIS observation for each orbit but it might also be a chance to estimate the fractional contribution of mosses across the landscape. Several studies found that satellite observations from multiple view angles

significantly improve the assessment of vertical vegetation structure<sup>38,39</sup>. At the same time, off-nadir view angles minimize distorting background/understory effects and provide new information for separating the spectral contribution of the forest floor (or understory layer) from the tree canopy layer<sup>40</sup>. Moreover, field manipulative experiments are needed to actually measure the contribution of bryophytes to satellite signals at the local scale, which will improve our understanding of the function of bryophytes.

Our current understanding of moss photosynthetic capability is still very limited. For example, the value of  $0.64 \,\mathrm{g C m^{-2} M J^{-1}}$ absorbed photosynthetically active radiation (APAR) for potential LUE of moss used in this study was 30% that of vascular plants according to currently available studies<sup>14</sup>. A previous study showed that the net  $\dot{CO}_2$  uptake rates of mosses were only about 10-40% those of vascular plants on a dry mass basis<sup>41</sup>. Similarly, the LUE of moss in Alaskan tundra was about 18% that of the overstory canopy<sup>42</sup>. Apparently, more comparative studies on photosynthetic capacity and LUE of mosses and vascular plants are needed to further improve our understanding of the role of mosses in regional and global carbon cycles. In addition, we lumped and did not differentiate the vascular plants (for example, shrubs and grasses) covering the ground of boreal forests from the trees in this study, as no systematic differences in photosynthetic capacity or LUE have been reported between these life forms. Nevertheless, understanding the importance of the contribution of ground vascular plants to the satellite signals in these forests versus their productivity and their contribution to ecosystem GPP can shed useful light to ecosystem carbon dynamics as the composition of these life forms changes over time<sup>18</sup>.

In this study, we found substantial impacts of disturbances on ecosystem structure and composition and their subsequent impacts on ecosystem productivity in boreal ecosystems. Disturbances alter forest structure and vegetation composition and cause higher bryophyte cover in young stands<sup>23</sup>. Increasing temperature in the northern high latitudes has resulted in welldocumented increases in frequency, intensity, seasonality and extent of various disturbances such as fire and insect outbreaks, and these trends are likely to continue with warming and drying<sup>43</sup>. The impacts of stand-replacing disturbances such as fire and harvesting might have been underestimated in the past, which can alter the forest age structure and contribution of moss to ecosystem GPP. It is imperative to incorporate these structural and functional changes of ecosystems into regional modelling frameworks to adequately assess the impacts of disturbances on the carbon cycle and global climate.

## Methods

**LUE model**. Predicting the GPP of terrestrial ecosystems has been a major challenge in quantifying the global carbon cycle<sup>44</sup>. Among predictive methods, the LUE model may have the greatest potential to adequately address the spatial and temporal dynamics of GPP because of its theoretical basis and practicality<sup>45</sup>. The LUE model is built on two fundamental assumptions<sup>46</sup>: (1) the ecosystem GPP is directly related to APAR through LUE, where LUE is defined as the amount of carbon produced per unit of APAR, and (2) realized LUE may be reduced below its theoretical potential value by environmental stresses, such as low temperatures or water shortages<sup>47</sup>. The general form of the LUE model is

$$GPP = fPAR \times PAR \times \varepsilon_{max} \times f \tag{1}$$

where fPAR is the fraction of PAR absorbed by the vegetation canopy, PAR is the incident photosynthetically active radiation (MJ m<sup>-2</sup>) per time period (for example, day or month),  $e_{max}$  is the potential LUE (g C m<sup>-2</sup> MJ<sup>-1</sup> APAR) without environmental stress, *f* is a scalar varying from 0 to 1 that represents the reduction of potential LUE under limiting environmental conditions, and the multiplication of  $e_{max}$  and *f* is realized LUE. Numerous studies have developed various LUE models such as the CASA

Numerous studies have developed various LUE models such as the CASA model<sup>16</sup>. In this study, three LUE models including MODIS-GPP, CASA and EC-LUE, which have been widely validated and applied at global scales, were selected

to examine the impact of moss on estimating vegetation GPP in the boreal forest biome.

**MODIS-GPP product.** The MODIS-GPP algorithm is an LUE-based model, with inputs from MODIS LAI/fPAR (MOD15A2), land cover and biome-specific climatological data sources from the National Aeronautics and Space Administration (NASA) Data Assimilation Office<sup>15</sup>. LUE is calculated based on two factors: the biome-specific potential LUE  $\varepsilon_{max}$ , a multiplier that reduces the LUE when cold temperatures limit plant function, and a second multiplier that reduces the potential LUE when vapour pressure deficit (VPD) is high enough to inhibit photosynthesis. It is assumed that soil water deficit covaries with VPD and that VPD will account for drought stress. The GPP algorithm was tested with flux data sets from a range of biomes<sup>48</sup>.

**CASA model.** CASA is a LUE model that uses satellite measurements to estimate vegetation NPP<sup>16</sup>. The model directly translates APAR into NPP based on LUE, which is a product of optimal efficiency and the regulatory functions of environmental factors (for example, temperature and water stress). The CASA model directly simulates NPP, and an approximate conversion of 0.47 between NPP and GPP is used in this study<sup>34</sup>.

**EC-LUE model**. The EC-LUE model is driven by only four variables: NDVI, PAR, air temperature (T, °C) and the evaporative fraction<sup>49</sup>. GPP was computed from fPAR, PAR and downregulators. The downregulators represent plant stresses due to suboptimal temperature  $(T_s)$  and water  $(W_s)$ .

$$GPP = fPAR \times PAR \times \varepsilon_{max} \times Min(T_s, W_s)$$
(2)

$$fPAR = 1.24 \times NDVI - 0.168$$
 (3)

$$T_{\rm s} = \frac{(T - T_{\rm min}) \times (T - T_{\rm max})}{((T - T_{\rm min}) \times (T - T_{\rm max}) - (T - T_{\rm opt})^2)}$$
(4)

$$W_{\rm s} = \frac{\rm LE}{R_{\rm n}} \tag{5}$$

where Min denotes the minimum values of  $T_s$  and  $W_s$ , and the impacts of temperature and moisture on LUE were assumed to follow Liebig's Law (that is, LUE is only affected by the most limiting factor at any given time).  $T_{\min}$ ,  $T_{\max}$  and  $T_{opt}$  are the minimum, maximum and optimum air temperatures (°C) for photosynthetic activity, respectively. If air temperature falls below  $T_{\min}$  or increases beyond  $T_{\max}$ ,  $T_s$  is set to zero. In this study,  $T_{\min}$  and  $T_{\max}$  were set to 0 and 40 °C, respectively, whereas  $T_{opt}$  was set to the mean temperature during the growing season. LE is the latent heat (MJ m<sup>-2</sup>), and MODIS-ET was the product used in this study.  $R_n$  is the net radiation (MJ m<sup>-2</sup>).

**EC-LUE-2p model.** To investigate the impact of moss on estimations of ecosystem GPP, we developed the EC-LUE-2p model by separating vegetation production into two components based on the EC-LUE model: vascular (GPP<sub>v</sub>) and moss (GPP<sub>m</sub>):

$$GPP = GPP_v + GPP_m \tag{6}$$

$$GPP_{v} = fPAR \times PAR_{v} \times (1 - K_{ND}VI) \times \varepsilon_{vmax} \times Min(W_{s}, T_{s})$$
(7)

$$GPP_{m} = fPAR \times PAR_{m} \times K_{ND} VI \times \varepsilon_{mmax} \times Min(W_{s}, T_{s})$$
(8)

where  $\varepsilon_{vmax}$  and  $\varepsilon_{mmax}$  are the potential LUE of vascular and moss plants, respectively. In this study,  $\varepsilon_{vmax}$  is set at 2.14 g C m<sup>-2</sup> MJ<sup>-1</sup> APAR<sup>12,50</sup>, and  $\varepsilon_{mmax}$  is assumed to be 30% of  $\varepsilon_{vmax}$  or 0.64 g C m<sup>-2</sup> MJ<sup>-1</sup> APAR based on previous studies<sup>14,41,42</sup>. K\_NDVI indicates the proportional contribution of moss to the NDVI signal; PAR<sub>v</sub> and PAR<sub>m</sub> are the total photosynthetically active radiation partitioned to the vascular and moss plants, respectively. The calculations of fPAR,  $W_s$  and  $T_s$  are reported in equations (3), (4) and (5).

We used the Beer-Lambert law to exponentially partition PAR between the vascular and moss surface:

$$PAR_{m} = PAR \times exp(-k \times LAI_{v})$$
(9)

$$PAR_{v} = PAR - PAR_{m}$$
(10)

$$LAI_v = (1 - K_NDVI) \times LAI$$
 (11)

where k is the light extinction coefficient (k = 0.5). LAI is MODIS LAI data, and the proportional contribution of vascular plants to the LAI signal is assumed to be the same as that of NDVI.

The linear relationship between NDVI and fPAR was used in EC-LUE, which was supported by many studies<sup>50,51</sup>. The strong linear relationship of NDVI and fPAR, derived from a total of 16 plant species of annuals, vines, deciduous and evergreen shrubs and trees, was used in our model<sup>52</sup>. Strong linear relationships between NDVI and field-measured fPAR were also observed at high-latitude ecosystems<sup>53,54</sup>. Moreover, we compared the MODIS-fPAR product and fPAR

calculations based on NDVI in this study, and results showed that they were comparable across almost all sites (Supplementary Fig. 6).

The EC-LUE-2p model and our study are built on two facts: first, ecosystem GPP includes contributions from higher plants and moss (K\_NDVI was used to represent contributions to the satellite signal from moss in the model). In case of no contribution from moss, K\_NDVI would be zero. Second, the intrinsic LUEs of higher plants and moss are different. This difference is the fundamental basis for our ability to separate the contributions of moss and vascular plants. If ecosystem GPP did not perform as highly as expected for vascular plants, it must contain mosses that have a substantially lower photosynthetic capacity. This is the biological foundation of the EC-LUE-2p model.

The only unknown in the model is K\_NDVI, which can be solved analytically according to equations (6)-(8):

$$K_NDVI = \frac{GPP - fPAR \times PAR_v \times \varepsilon_{vmax} \times Min(W_s, T_s)}{fPAR \times PAR_w \times \varepsilon_{mmax} \times Min(W_s, T_s) - fPAR \times PAR_v \times \varepsilon_{vmax} \times Min(W_s, T_s)}$$
(12)

In reality, K\_NDVI is solved statistically using estimates of GPP from EC measurements, as we have done in this study. The biological and physiological bases of the model (two facts described above) and the statistical optimization of K\_NDVI against EC measurements strictly and uniquely constrained the meaning of K\_NDVI.

For the EC-LUE-2p, PAR,  $R_n$ , T and LAI are model forcing data that can be measured by EC towers and derived from satellite data. Potential LUEs of vascular and moss plants (that is,  $v_{vmax}$  and  $v_{mmax}$ ) are model parameters, which have been determined as mentioned above. The only undetermined parameter is K\_NDVI, which is optimized at each EC site against GPP estimates. The regression procedure (Proc NLIN) in the Statistical Analysis System (SAS, SAS Institute, Cary, NC, USA; any use of trade, firm or product names is for descriptive purposes only and does not imply endorsement by the US Government) was applied to optimize the values for K\_NDVI at each site.

At the regional scale, the relationship between K\_NDVI and stand age (Fig. 3) was used to calculate the spatial pattern of K\_NDVI, based on the spatial distribution map of forest age in Canada<sup>25,55</sup>. We validated this large-scale model application at the site level. First, we used the relationship shown in Fig. 3 for each site individually and calculated K\_NDVI to simulate GPP. Second, the correlation relationship between K\_NDVI and stand age was recalculated by excluding a given site, and the relationship was used to estimate K\_NDVI and GPP at this site. Simulations of GPP at these two schemes were then compared with observations. The results showed good model performance using the relationship between K\_NDVI and stand age (Supplementary Fig. 7).

**Model operation and forcing data at the EC sites.** The EC-LUE and CASA models were run at 8-day time steps at the EC sites. Environmental variables measured at the EC sites were used to drive these two models. The 8-day NDVI data with 1-km spatial resolution (MOD13) were retrieved at the EC sites. MOD13 has a quality control (QC) layer to indicate data reliability for each pixel (good data, marginal data, snow/ice, cloudy and missing). In this study, we considered NDVI data with a 'marginal' or 'cloudy' QC flag to be unreliable. We temporally filled the missing or unreliable NDVI at each 1-km MODIS pixel based on its corresponding quality assessment data field<sup>56</sup>. First, if the first (or last) 8-day NDVI value is unreliable or missing, it will be replaced by the closest reliable 8-day value. Second, other unreliable NDVI will be replaced by linear interpolation of the nearest reliable value before it and the closest reliable value after it. If there are no reliable NDVI values during the entire year, the annual maximum NDVI will be chosen from unreliable periods in the current year and will be used as a constant value across the entire year.

The 8-day MODIS evapotranspiration (ET) data (MOD16)<sup>57</sup> at 1-km spatial resolution was used to indicate actual ET, and potential ET (PET) was calculated by the method of Thornthwaite<sup>58</sup>. The 8-day MODIS-GPP product (MOD17) data at 1-km spatial resolution were used to examine MODIS-GPP algorithm verification at the EC sites. The same method that was used for NDVI was used to conduct data QC and fill missing data gaps for MODIS-ET and -GPP products.

**Statistical analysis.** Three metrics were used to evaluate the performance of the four models in this study: the coefficient of determination ( $R^2$ ), absolute predictive error (PE) and RPE. PE quantified the difference between simulated and observed values, and RPE was computed as

$$RPE = \frac{\overline{S} - \overline{O}}{\overline{O}} \times 100 \%$$
(13)

where  $\overline{S}$  and  $\overline{O}$  are mean simulated and mean observed values, respectively.

**Field-measured LAI at seven Canada sites.** Field-measured LAI of vascular plants at seven adjacent Canada sites recovering from fire disturbances (that is, CA-NS1, CA-NS2, CA-NS3, CA-NS5, CA-NS95 and CA-NS7) was used to compare with satellite NDVI values. LAI was measured by harvest<sup>20</sup>. The harvests occurred shortly before the end of the growing season in 2004. Moreover, we used 8-day MODIS-NDVI product (MOD13). Figure 2 shows the mean NDVI values from May to September during 2002–2006, and the error bars indicate s.d.

**Model forcing data over the regional scale**. We used input data sets of  $R_{n}$ , T and PAR from the Modern Era Retrospective-Analysis for Research and Applications (MERRA)<sup>59</sup> archive in 2004 to drive the CASA and EC-LUE models for regional GPP estimation. The MODIS-GPP (MOD17) product was directly used in this study. MERRA is a NASA reanalysis for the satellite era using a major new version

of the Goddard Earth Observing System Data Assimilation System Version 5 (GEOS-5), which was used to interpolate and grid these point data at short time intervals, producing an estimate of climatic conditions for the world at 10 m above the land surface (approximating canopy height conditions) and at a resolution of  $0.5^{\circ}$  latitude by  $0.6^{\circ}$  longitude. The MERRA reanalysis data set has been validated carefully at the global scale using surface meteorological data sets to evaluate the uncertainty of various meteorological factors (for example, temperature, radiation, humidity and energy balance)<sup>60</sup>. This validation showed that MERRA considerably reduced imbalances in energy and water<sup>61</sup>. Detailed information on the MERRA data set is available at http://gmao.gsfc.nasa.gov/research/merra/.

Eight-day MODIS LAI/fPAR (MOD15), NDVI (MOD13) and ET (MOD16) at 1-km spatial resolution also were used to drive the CASA, EC-LUE and EC-LUE-2p models for estimating regional GPP. The same method as that used at the EC sites was used to conduct data QC and fill missing data gaps for MODIS products. All satellite-based products were resampled at a resolution of 0.5° latitude by 0.6° longitude to match that of the MERRA reanalysis data set.

#### Uncertainty of GPP Estimations. Figure 3 shows the correlation between

K\_NDVI and stand age. To quantify the uncertainty, we calculated the 95% confidence limits of the regression relationship between stand age and K\_NDVI. The high and low 95% confidence limits were used to quantify the maximum and minimum GPP estimations, respectively.

This study used MODIS NDVI and LAI products of 1-km resolution to estimate regional GPP using the EC-LUE, EC-LUE-2p and CASA models. These products were generated from MODIS surface reflectance after correcting for the effects of cloud contamination, ozone absorption and aerosols, and adjusted to nadir and standard sun angles using bidirectional reflectance distribution function models<sup>62</sup>. A simple linear interpolation method was used to fill unreliable or missing NDVI or LAI data<sup>56</sup>, which may introduce additional uncertainty into our GPP estimation. Moreover, we used the relationship between stand age and contribution of moss to the satellite signal to improve regional GPP estimations. However, we did not consider impacts of other factors on moss fraction cover, such as vegetation, soils and surface water. On the basis of long-term measurement and literature data, synthesis analysis showed substantial differences of moss cover in the uplands and floodplains<sup>63</sup>.

## References

- Aaron, R. & Gibbs, H. K. New IPCC Tier-1 global biomass carbon map for the year 2000, http://cdiac.ornl.gov (2008).
- Tarnocai, C. et al. Soil organic carbon pools in the northern circumpolar permafrost region. Glob. Biogeochem. Cycles 23, GB2023 (2009).
- 3. Pan, Y. D. *et al.* A large and persistent carbon sink in the World's forests. *Science* **333**, 988–993 (2011).
- IPCC. Climate Change 2007: the Physical Science Basis (Cambridge Univ. Press, 2007).
- Schaphoff, S. et al. Contribution of permafrost soils to the global carbon budget. Environ. Res. Lett. 8, 014026 (2013).
- Chapin, F. S., Matson, P. A., Vitousek, P. & Chapin, M. C. Principles of Terrestrial Ecosystem Ecology (Springer, 2012).
- Cramer, W. *et al.* Comparing global models of terrestrial net primary productivity (NPP): overview and key results. *Glob. Change Biol.* 5, 1–15 (1999).
- Huntzinger, D. N. *et al.* North American Carbon Program (NACP) regional interim synthesis: terrestrial biospheric model intercomparison. *Ecol. Model.* 232, 144–157 (2012).
- Liu, J., Chen, J. M., Cihlar, J. & Chen, W. Net primary productivity mapped for Canada at 1-km resolution. *Global Ecol. Biogeogr.* 11, 115–129 (2002).
- Kimball, J. S., Zhao, M. S., McDonald, K. C. & Running, S. W. Satellite remote sensing of terrestrial net primary production for the pan-arctic basin and Alaska. *Mitig. Adapt. Strategies Glob. Change* 11, 783–804 (2006).
- 11. Stinson, G. et al. An inventory-based analysis of Canada's managed forest carbon dynamics, 1990 to 2008. Glob. Change Biol. 17, 2227–2244 (2011).
- Yuan, W. P. et al. Global estimates of evapotranspiration and gross primary production based on MODIS and global meteorology data. *Remote Sens. Environ.* 114, 1416–1431 (2010).
- Peckham, S. D., Ahl, D. E. & Gower, S. T. Bryophyte cover estimation in a boreal black spruce forest using airborne lidar and multispectral sensors. *Remote Sens. Environ.* 113, 1127–1132 (2009).
- Whitehead, D. & Gower, S. T. Photosynthesis and light-use efficiency by plants in a Canadian boreal forest ecosystem. *Tree Physiol.* 21, 925–929 (2001).
- 15. Running, S. W., Nemani, R., Glassy, J. M. & Thornton, P. E. MODIS daily photosynthesis (PSN) and annual net primary production (NPP)

product (MOD17), algorithm theoretical basis document, version 3.0. http://modis.gsfc.nasa.gov/ (1999).

- Potter, C. S. et al. Terrestrial ecosystem production: a process model based on global satellite and surface data. *Glob. Biogeochem. Cycles* 7, 811–841 (1993).
- Schimmel, J. & Granstrom, A. Fire severity and vegetation response in the boreal Swedish forest. *Ecology* 77, 1436–1450 (1996).
- Liu, H. P. & Randerson, J. T. Interannual variability of surface energy exchange depends on stand age in a boreal forest fire chronosequence. *J. Geophys. Res.* 113, G01006 (2008).
- Kuusk, A., Lang, M. & Nilson, T. Simulation of the reflectance of ground vegetation in sub-boreal forests. Agric. For. Meteorol. 126, 33–46 (2004).
- Goulden, M. L. et al. Patterns of NPP, GPP, respiration, and NEP during boreal forest succession. Glob. Change Biol. 17, 855–871 (2011).
- Masek, J. G. Stability of boreal forest stands during recent climate change: evidence from Landsat satellite imagery. J. Biogeogr. 28, 967–976 (2001).
- Frolking, S. Sensitivity of spruce/moss boreal forest net ecosystem productivity to seasonal anomalies in weather. J. Geophys. Res. 102, 29053–29064 (1997).
- Hodge, D. A., Pharo, E. J., Dalton, P. J. & Turner, P. A. M. Successional patterns of terrestrial bryophytes along a wildfire chronosequence in the wet eucalypt forests of southern Tasmania. *Tasforests* 18, 67–76 (2009).
- Turetsky, M. R., Mack, M. C., Hollingsworth, T. N. & Harden, J. W. The role of mosses in ecosystem succession and function in Alaska's boreal forest. *Can. J. For. Res.* 40, 1237–1264 (2010).
- Chen, J. M. et al. Spatial distribution of carbon sources and sinks in Canada's forests. Tellus B 55, 622–641 (2003).
- Beer, C. et al. Terrestrial gross carbon dioxide uptake: global distribution and covariation with climate. Science 329, 834–838 (2010).
- Mahadevan, P. *et al.* A satellite-based biosphere parameterization for net ecosystem CO<sub>2</sub> exchange: vegetation photosynthesis and respiration model (VPRM). *Glob. Biogeochem. Cycles* 22, 1–17 (2008).
- Deng, F. et al. Global monthly CO<sub>2</sub> flux inversion with a focus over North America. Tellus B 59, 179–190 (2007).
- 29. Dargaville, R., McGuire, A. D. & Rayner, P. Estimates of large scale fluxes in high latitudes from terrestrial biosphere models and an inversion of atmospheric CO<sub>2</sub> measurements. *Clim. Change* **55**, 273–285 (2002).
- 30. McGuire, A. D. *et al.* Sensitivity of the carbon cycle in the Arctic to climate change. *Ecol. Monogr.* **79**, 523–555 (2009).
- Hayes, D. J. et al. Is the northern high-latitude land-based CO2 sink weakening? Glob. Biogeochem. Cycles 25, GB3018 (2011).
- Proctor, M. C. F. in *Photosynthesis in Bryophytes and Early Land Plants* Vol. 37, Advances in Photosynthesis and Respiration (eds Hanson, D. T. & Rice, S. K.) 59–77 (Springer Science and Business Media, Dordrecht, 2014).
- 33. Ryu, Y. *et al.* Integration of MODIS land and atmosphere products with a coupled-process model to estimate gross primary productivity and evapotranspiration from 1 km to global scales. *Glob. Biogeochem. Cycles* **25**, GB4017 (2011).
- Waring, H., Landsberg, J. J. & Williams, M. Net primary production of forests: a constant fraction of gross primary production? *Tree Physiol.* 18, 129–134 (1998).
- Street, L. E., Subke, J. A., Sommerkorn, M., Heinemeyer, A. & Williams, M. Turnover of recently assimilated carbon in arctic bryophytes. *Oecologia* 167, 325–337 (2011).
- Gower, S. T. et al. Net primary production and carbon allocation patterns of boreal forest ecosystems. Ecol. Appl. 11, 1395–1411 (2001).
- Bond-Lamberty, B. & Gower, S. T. Estimation of stand-level leaf area for boreal bryophytes. *Oecologia* 151, 584–592 (2007).
- Vuolo, F., Dini, L. & D'Urso, G. Retrieval of leaf area index from CHRIS/ PROBA data: an analysis of the directional and spectral information content. *Int. J. Remote Sens.* 29, 5063–5072 (2008).
- Chen, J. M., Li, X., Nilson, T. & Strahler, A. Recent advances in geometrical optical modelling and its applications. *Remote Sens. Rev.* 18, 227–262 (2000).
- Rautiainen, M. *et al.* Multi-angular reflectance properties of a hemiboreal forest: an analysis using CHRIS PROBA data. *Remote Sens. Environ.* 112, 2627–2642 (2008).
- Martin, C. E. & Adamson, V. J. Photosynthetic capacity of mosses relative to vascular plants. J. Bryol. 23, 319–323 (2001).
- Huemmrich, K. F. et al. Remote sensing of tundra gross ecosystems productivity and light use efficiency under varying temperature and moisture conditions. *Remote Sens. Environ.* 114, 481–489 (2010).
- 43. Kasischke, E. S. & Turetsky, M. R. Recent changes in the fire regime across the North American boreal region—spatial and temporal patterns of burning across Canada and Alaska. *Geophys. Res. Lett.* 33, L09703 (2006).
- Canadell, J. G., Steffen, W. L. & White, P. S. Carbon metabolism of the terrestrial biosphere: a multi-technique approach for improved understanding. *Ecosystems* 3, 115–130 (2000).
- Monteith, J. L. Climate and the efficiency of crop production in Britain. *Phil. Trans. R. Soc. Lond. B* 281, 277–294 (1977).

# ARTICLE

- Running, S. W. *et al.* A continuous satellite-derived measure of global terrestrial primary production. *Bioscience* 54, 547–560 (2004).
- Landsberg, J. J. Physiological Ecology of Forest Production 165–178 (Academic Press, 1986).
- Heinsch, F. A. Evaluation of remote sensing based terrestrial production from MODIS using AmeriFlux eddy tower flux network observations. *IEEE Trans. Geosci. Remote Sensing* 44, 1908–1925 (2006).
- 49. Yuan, W. P. et al. Deriving a light use efficiency model from eddy covariance flux data for predicting daily gross primary production across biomes. Agric. For. Meteorol. 143, 189–207 (2007).
- Myneni, R. B., Hall, F. G., Sellers, P. J. & Marshak, A. L. The meaning of spectral vegetation indices. *IEEE Trans. Geosci. Remote Sensing* 33, 481–486 (1995).
- Myneni, R. B. *et al.* Global products of vegetation leaf area and fraction absorbed PAR from year one of MODIS data. *Remote Sens. Environ.* 83, 214–231 (2002).
- 52. Sims, D. A. *et al.* Midday values of gross CO<sub>2</sub> flux and light use efficiency during satellite overpasses can be used to directly estimate eight-day mean flux. *Agric. For. Meteorol.* 131, 1-12 (2005).
- Olofsson, P. & Eklundh, L. Estimation of absorbed PAR across Scandinavia from satellite measurements. Part II: modeling and evaluating the fractional absorption. *Remote Sens. Environ.* 110, 240–251 (2007).
- Steinberg, D. C. & Goetz, S. Assessment and extension of the MODIS FPAR products in temperate forests of the eastern United States. *Int. J. Remote Sens.* 30, 169–187 (2009).
- 55. Pan, Y. D. *et al.* Age structure and disturbance legacy of North American forests. *Biogeosciences* **8**, 715–732 (2011).
- Zhao, M. S., Heinsch, F. A., Nemani, R. R. & Running, S. W. Improvements of the MODIS terrestrial gross and net primary production global data set. *Remote Sens. Environ.* 95, 164–176 (2005).
- Mu, Q. Z., Zhao, M. S. & Running, S. W. Improvements to a MODIS global terrestrial evapotranspiration algorithm. *Remote Sens. Environ.* 115, 1781–1800 (2011).
- Thornthwaite, C. W. An approach toward a rational classification of climate. Geogr. Rev. 38, 55-94 (1948).
- Global Modeling and Assimilation Office. File Specification for GEOSDAS Gridded Output Version 5.3, Report (NASA Goddard Space Flight Center, Greenbelt, MD, 2004).
- Rienecker, M. M. et al. MERRA: NASA's Modern-Era Retrospective analysis for research and applications. J. Climate 24, 3624–3648 (2011).
- 61. Kennedy, A. D. et al. A comparison of MERRA and NARR reanalyses with the DOE ARM SGP data. J. Climate 24, 4541-4557 (2011).

- Huete, A., Justice, C. & van Leeuwen, W. MODIS vegetation index (MOD13) algorightm theoretical basis document. University of Arizona Version 3, http:// modis.gsfc.nasa.gov/data/atbd/atbd\_mod13.pdf (1999).
- Rapalee, G., Steyaert, L. T. & Hall, F. G. Moss and lichen cover mapping at local and regional scales in the boreal forest ecosystem of central Canada. J. Geophys. Res. Atmos. 106, 33551–33563 (2001).

### Acknowledgements

This study was supported by the National Key Program for Global Change Research of China (2010CB950504), National Science Foundation for Excellent Young Scholars of China (41322005), Science Fund for Creative Research Groups of the National Natural Science Foundation of China (41121001) and Freedom Project (No. SKLCS-ZZ-2013-02-02) of the State Key Laboratory of Cryospheric Sciences, Cold and Arid Regions Environmental and Engineering Research Institute, Chinese Academy of Sciences and Program for New Century Excellent Talents in University (NCET-12-0060). The research was also partially funded by the Geographic Analysis and Monitoring (GAM) Program and the Global Change Research and Development Program (R&D) of the US Geological Survey. We acknowledge the work of principal investigators and collaborators in EC sites, who provided the EC flux measurements. The following networks participated with flux data: Fluxnet-Canada, AmeriFlux, CarboEuropeIP, Asiaflux and TCOS-Siberia network.

#### Author contributions

S.Liu., W.Y., W.D. and S.Lia. designed the study. W.Y. and S.Liu. performed the analyses. W.Y., S.Liu., S.Z. and W.D. drafted the paper. S.Liu., J.C., E.M., T.V. and L.K. contributed to the interpretation of the results. A.B., T.A.B., M.L.G., L.K., A.L., H.A.M., Y.M., E.M., M.M., T.O., K.P., A.V., T.M. contributed the eddy covariance flux measurements. All authors contributed extensively to the revisions of the manuscript.

#### Additional information

Supplementary Information accompanies this paper at http://www.nature.com/ naturecommunications

Competing financial interests: The authors declare no competing financial interests.

**Reprints and permission** information is available online at http://npg.nature.com/ reprintsandpermissions/

How to cite this article: Yuan, W. *et al.* Differentiating moss from higher plants is critical in studying the carbon cycle of the boreal biome. *Nat. Commun.* 5:4270 doi: 10.1038/ncomms5270 (2014).